

ORIGINAL RESEARCH

Open Access



Restoring degraded agricultural peatlands: how rewetting, biochar, and iron sulphate synergistically modify microbial hotspots and carbon storage

Peduruhewa H. Jeewani^{1*}, Robert W. Brown¹, Jennifer M. Rhymes², Chris D. Evans², Dave R. Chadwick¹ and Davey L. Jones¹

Abstract

The draining and conversion of peatlands for agriculture has led to their degradation globally, diminishing their carbon (C) storage capacity and functioning. However, rewetting, alongside the addition of organic/inorganic amendments, has the potential to accelerate peat formation and C accrual. The aim of this experiment was therefore to examine the combined benefits of altering water table depth and adding organic (e.g., biochar, paper waste, biosolids, cereal straw; 20 t C ha⁻¹) and inorganic (e.g., FeSO₄; 0.5 t ha⁻¹) materials on net C storage and peatland functioning (i.e., microbial communities, greenhouse gas emissions and biogeochemical cycling). The experiment consisted of outdoor agricultural peat mesocosms monitored over 1 year. The relative effectiveness of the amendments in preserving peat-C (t C ha⁻¹) followed the series: *Miscanthus* biochar (18.9 t C ha⁻¹) > *Miscanthus* residues (17.3 t C ha⁻¹) > biosolids (17.2 t C ha⁻¹) > cereal straw (14.5 t C ha⁻¹) > paper waste (13.3 t C ha⁻¹) based on C additional rate (20 t C ha⁻¹). Overall, a high-water table combined with biochar and FeSO₄ addition was the most effective at suppressing enzyme activity (e.g., β-glucosidase, phenol oxidase, cellobiase), methanogen activity (e.g., *Methanosarcina*) and peat mineralization rate. We ascribe this in part to changes in the fungal and bacterial community structure (e.g., reductions in Actinobacteria by – 22% and Ascomycota by – 61%). FeSO₄ also increased the Fe-bound C content in the non-rewetted treatment, supporting the ‘iron gate’ mechanism for C preservation. The mechanisms behind our results appear to be both abiotic (affecting SOC solubility through changes in redox conditions and Fe–C interactions) and biotic (via shifts in microbial community and enzyme activities), creating conditions that enhance C preservation. These findings provide evidence for implementing biochar and FeSO₄ amendments alongside water table management as practical, scalable strategies for restoring C storage capacity in agricultural peatlands.

Highlights

- Rewetting with biochar and FeSO₄ suppresses decomposers and enhances peat C storage.
- FeSO₄ suppresses methanogens and increases Fe-protected C in aerobic conditions.
- Biochar modifies the microbial community and reduces hydrolytic enzyme activity.
- Strategic amendment combinations reduce GHG emissions while preserving peat soil C.

*Correspondence:

Peduruhewa H. Jeewani
j.hemamali@bangor.ac.uk

Full list of author information is available at the end of the article



© Crown 2025. **Open Access** This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing, adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons licence, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons licence, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons licence and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this licence, visit <http://creativecommons.org/licenses/by/4.0/>.

- Raising water tables prevents dominance of Actinobacteria and Ascomycota decomposers.

Keywords Soil solution, Hydrolytic enzymes, Sustainable agriculture, Soil microbes, Fe gate, Histosol

1 Introduction

Peatlands account for approximately 2.84% of the global land area, constituting the largest terrestrial carbon (C) reservoir (Evans et al. 2017; Evans et al. 2023). Despite their critical role in carbon storage, an estimated 20% of the world's peatlands are currently under agricultural use, which significantly alters their natural carbon dynamics and contributes to greenhouse gas emissions (Bonn et al. 2016; Xu et al. 2018; Kwon et al. 2022). Drainage of lowland peat soils is a common method of increasing agricultural productivity, leading to significant degradation and loss of C through mineralization (Andersen et al. 2013; Brouns et al. 2016). Rewetting of peat soils and application of soil amendments (for example, biochar; the product of pyrolysis of organic feedstock) are common restoration practices (Evans et al. 2023; Deshoux et al. 2023; Jeewani et al. 2025), yet our understanding of the belowground biogeochemistry, particularly C and nitrogen (N) cycling, is still incomplete. Thus, it is important to consider how peat soil biogeochemistry is affected by these new practices, particularly when estimating the current and future responses of peatland C balance.

Peatlands contain a diverse fungal community with significant metabolic diversity that play a key role in organic matter transformation (Andersen et al. 2013). Among these fungi, Ascomycota is believed to be the most abundant phylum responsible for aerobic decomposition in peatlands (Pankratov et al. 2006). Peltoniemi et al. (2009) observed that the peat fungal community varied between sites representing the prevailing nutrient and water table regime. However, contrasting results have been observed among the peat fungal community in response to lowering of the water table (Peltoniemi et al. 2009; Wang et al. 2024b). Changes in the composition and size of the bacterial community have also been observed in response to changes in water table level (Deshoux et al. 2023). Typically, the community becomes less diverse when the water table is raised, a change which has been ascribed to the imposition of anaerobic conditions, decreased competition and increased rates of predation (Andersen et al. 2013). In peatlands, the boundary zone between the oxic and anoxic layers, where CH₄ and O₂ are both present, also creates an ideal niche for methanotrophs, the most prevalent bacterial phyla of which are *Proteobacteria* (*gamma*- and *alpha*-*Proteobacteria*) and *Verrucomicrobia* (Op den Camp et al. 2009). Given the diverse microbial responses to environmental

conditions in peatlands, it is essential to investigate how microbial community structure and function change in response to management interventions such as rewetting and amendment addition, particularly regarding their effects on C storage and greenhouse gas (GHG) emissions.

Incorporation or surface broadcasting of organic amendments re-introduces C that may have previously been lost due to peatland drainage and cultivation. However, the success in restoring C is likely regulated by the biogeochemistry of the added substrate, moisture status and their combined impact on the composition and functioning of the soil microbial community (Eilers et al. 2010; Butterbach-Bahl et al. 2013; Kandel et al. 2019; Brown et al. 2022). For example, due to its aromatic structure, biochar-derived C is highly persistent in soil, while other C substrates (e.g., straw) may be less persistent and generate unwanted CH₄ emissions (Liu et al. 2014; Wang et al. 2024b). To date, however, our mechanistic understanding of the fate and behavior of different C substrates in rewetted agricultural peatlands remains poor.

Peat soil biogeochemistry can be strongly influenced by the presence of alternative electron acceptors (e.g., Fe-oxyhydroxides, NO₃⁻, SO₄²⁻), suggesting that they have the potential to be used to manipulate microbial functioning during peatland restoration (Dean et al. 2018; Richy et al. 2024). For example, incubation experiments have observed that the addition of Fe³⁺ and SO₄²⁻ may limit both denitrification and methanogenesis (Wen et al. 2019; Jeewani et al. 2025). The presence of SO₄²⁻ may also promote the growth of SO₄²⁻-reducing bacteria such as *Desulfovibrio spp.*, *Desulfosarcina spp.* and *Desulfobacter spp.* and archaea (*Archaeoglobus spp.*), leading to increased competition and a suppression of methanogen activity (Pester et al. 2012; Ozuolmez et al. 2015). Furthermore, earlier field studies have demonstrated that rewetting iron-rich peat soils can still lead to substantial CH₄ emissions, even though methanogenesis is expected to be suppressed by the presence of iron oxides (Yamada et al. 2014). However, the impact of FeSO₄ addition on the GHG release and net C storage in agricultural peat soils remains poorly understood.

To the best of our knowledge, this is the first study to investigate how the combined application of inorganic and organic amendments, along with water table adjustments, influences microbial community dynamics in

agricultural peatland soils. In this study, our goal was to identify the most effective amendments for enhancing soil carbon content without promoting the decomposition of soil organic matter (SOM) or increasing net GHG emissions. We hypothesized that: 1) Under conditions of elevated water table, the addition of high C:N ratio amendments (e.g., *Miscanthus* biochar) will limit nitrogen availability to decomposers, thereby constraining microbial growth and enzyme-mediated decomposition processes via nutrient immobilization, relative to low C:N treatments; 2) Low C:N organic amendments (e.g., biosolids, paper waste) will stimulate rapid microbial turnover and SOM decomposition by promoting the growth of copiotrophs (r-strategists), leading to increased enzyme activity and CO₂ fluxes; 3) FeSO₄ addition will suppress methanogenesis by favoring alternative terminal electron-accepting processes (e.g., iron reduction), leading to shifts in the abundance of methanogenic archaea and lower CH₄ emissions.

2 Materials and methods

2.1 Site information

Intact soil cores were obtained from a commercially managed lowland agricultural peat field in Nottinghamshire, UK (53°27'N, 00°54'W) in March 2022. The site is a level, drained lowland fen characterized by an approximately 0.5-m-thick organic layer resting atop mineral soil and had been under intensive agricultural production for over 100 years. The average annual temperature of the site is 10.3 °C and annual rainfall of 1162 mm, and the soil is classified as an Ombric Sapric Histosols (Mantel, et al., 2023)). Intact soil cores ($n=56$; depth=50 cm) were collected using PVC pipes ($\varnothing=20$ cm, height=60 cm). The intact soil cores were transported to Bangor University, where they were kept outdoors throughout the entire 365-day measurement period.

2.2 Basic properties of organic amendments and soil

Initial soil (0–10 cm depth) and organic amendment characteristics were quantified at the start of the experiment. The pH and electrical conductivity (EC) of soil samples were tested in 1:2.5 (w/v) soil-to-distilled water suspensions using a standard probe. Bulk density was measured using the fixed-volume ring technique, following the procedure detailed by Blake and Hartge (1986). Total carbon (TC) and total nitrogen (TN) levels in both soil samples and organic amendments were analyzed using oven-dried (80 °C for 24 h), finely ground material on a TruSpec[®] CN Analyzer (Leco Corp., St. Joseph, MI). SOM content was assessed through the loss-on-ignition method, conducted in a muffle Furnace at 450 °C for a duration of 16 h. Dissolved organic carbon (DOC) and total dissolved nitrogen (TDN) were quantified from a

0.5 M potassium sulfate (K₂SO₄) extraction using a Multi N/C 2100/2100 analyzer (Analytik Jena AG, Jena, Germany). Nutrient concentrations including ammonium (NH₄⁺), nitrate (NO₃⁻), phosphate (PO₄³⁻), and sulfate (SO₄²⁻) in the amendments were extracted with distilled water at a 1:5 (w/v) ratio and analyzed spectrophotometrically using a Power Wave-Xs microplate reader (BioTek Instruments Inc., Winooski, VT, USA), based on established colorimetric protocols (Murphy & Riley, 1962; Bradfield & Cooke, 1985; Kuo & Sparks, 1996). Microbial biomass carbon (MBC) was estimated through the chloroform fumigation–extraction technique, as described by Brookes et al. (1985) and Vance et al. (1987). The biochar material was characterized by determining its atomic hydrogen-to-carbon (H/C) ratio, and the stable polyaromatic carbon (SPAC) content was measured via hydro-pyrolysis (Hypy) following the methodology of Ascough et al. (2009). The detailed characteristics of the soil and organic amendments are provided in Table S1 (Jeewani et al. 2025).

2.3 Experimental setup

The mesocosm experiment comprised 14 treatments (each with four replicates), incorporating five different organic amendments that varied in their C:N ratios. The treatments were: (i) Control soil cores with a water table (WT) at – 40 cm (Control+LW; to represent no change of management), (ii) Control soil cores with a saturated water table at the soil surface (0 cm; Control+HW), (iii) high WT + pyrolysed *Miscanthus giganteus* wood chip (biochar; pyrolysed at 450 °C, 30 min; C:N ratio=258; Biochar+HW), (iv) high WT + commercial paper waste (C:N ratio=155; Paperwaste+HW), (v) high WT + *M. giganteus*-derived chip (C:N ratio=96; M.chip+HW), (vi) high WT + Cereal straw (*Hordeum vulgare* L.) (C:N ratio=63; C.straw+HW), and (vii) high WT + anaerobically digested biosolids sourced from a high-capacity urban wastewater treatment plant (C:N ratio=10; Biosolids+HW). Each treatment was assessed with or without the addition of FeSO₄ (Table S2; Fig.S3). All organic substrates were applied at a rate of 20 t C ha⁻¹ to simulate field-level organic residue management practices (Jones et al. 2012; Pandit et al. 2018). while FeSO₄ was added at in accordance with Wen et al. (2019). The FeSO₄·7H₂O addition rate (0.5 t ha⁻¹) was selected based on prior peatland amendment studies aiming to stimulate Fe–C complexation and suppress methanogenesis without exceeding levels known to cause toxicity or alter redox buffering capacity (Wen et al. 2019; Hu et al. 2020). Both inorganic (FeSO₄) and organic amendments were manually incorporated into the upper 10 cm of the soil to emulate mechanical harrowing under field conditions.

Each mesocosm was positioned within a larger external container filled with water, functioning as a bund system. Drainage holes were installed to regulate the water table at either a high level (0 cm) or a low level (−40 cm). Prior to initiating the experiment, mesocosms were stabilized at their designated water table depths for a period of three days. Throughout the experimental period, water levels were sustained via natural rainfall, supplemented with tap water during dry intervals. The selected water table settings were intended to establish primarily anaerobic conditions in the high-water table (HW) or 're-wetted' treatments, and predominantly aerobic conditions in the low water table (LW) control, simulating conventional drainage-based agricultural practices.

2.4 Soil solution chemistry

Horizontally installed rhizon samplers (Rhizosphere Research Products, Wageningen, The Netherlands) were used to collect soil solution from soil depths of 5 cm (in the treatment layer) and 25 cm (below the treatment layer). Samples were collected on days 1, 3, 5 and 7 then weekly and monthly, parallel to the GHG sample collections (Jeewani et al. 2025). On collection, the samples were stored at (−20 °C for <14 days) until analysis for TDN, DOC, NH_4^+ , NO_3^- , PO_4^{3-} and SO_4^{2-} , and were quantified using the methods described in Sect. 2.2. At each sampling event, in situ measurements of soil pH, temperature (T), redox potential (Eh), and electrical conductivity (EC) were conducted within the upper 5 cm of the soil profile using a Sension+MM150 portable sensor system (Hach UK, Manchester, UK). Additionally, the DOM (filtered through 0.45 μm) was measured by specific UV absorbance at 254 nm using a Cary 60 UV–vis spectrophotometer (UV–Vis, Agilent Technologies, UK). This has been shown to be indicative of the hydrophobic organic acid fraction, as well as aromatic content and molecular weight of compounds present (Spencer et al. 2012; Chowdhury 2013).

2.5 Fe-bound organic C determination

Fe-organic carbon complexes were quantified using the dithionite citrate bicarbonate (DCB) method after 180 days, following Wang et al. (2017). Briefly, lyophilized soil (0–10 cm depth) was mixed with a buffer solution at a 1:60 (w/v) ratio: a solution containing 0.27 M trisodium citrate and 0.11 M sodium bicarbonate, adjusted to pH 7.3 and heated to 80 °C. Then, 0.5 g of sodium dithionite was added as a reducing agent, and the mixture was maintained at 60–80 °C for 15 min. For the control treatment, the soil samples were extracted with NaCl at an equivalent ionic strength instead of undergoing DCB extraction. The residues were subsequently rinsed three times with 1 M NaCl, then were oven-dried for

SOC analysis, following the procedure outlined previously. The concentration of Fe-bound organic carbon was determined by calculating the difference in SOC content between the control and the reduction treatments. Soil Fe(II) and Fe(III) contents were analyzed using the ferrozine-ultraviolet absorbance method (Wang et al. 2017). Specifically, fresh soil (0–10 cm depth) was extracted with 0.5 M HCl overnight (1:25 w/v). Fe(II) concentration was quantified by measuring absorbance at 562 nm with a UV–Vis spectrophotometer, following reaction with a 5 mM ferrozine solution (Stokey, 1970). Total iron was quantified after reduction with 2% hydroxylamine hydrochloride, with Fe(II) subsequently measured as described above. The amount of Fe(III) was calculated by subtracting the Fe(II) value from the total iron and expressed as mg g^{-1} soil.

2.6 Extracellular enzyme activity

Enzyme activity kinetics were evaluated using fluorogenic 4-methylumbelliferone (MUF) substrates (Razavi et al. 2015). The MUF substrates were diluted with sterile MES buffer to the required concentrations, and enzyme activities were measured across a gradient of substrate concentrations (0, 10, 20, 30, 40, 50, 100, 200 $\mu\text{mol g}^{-1}$ soil). Briefly, soil suspensions (1:50, w/v; soil-to-deionized H_2O) were shaken for 2 min (Koch et al. 2007). Subsequently, 50 μL of the soil suspension was combined with 50 μL of buffer (pH 6.5) and 100 μL of the appropriate substrate solution in a 96-well microplate (Puregrade, Germany). Calibration solutions were prepared by combining 50 μL of the soil suspension with varying concentrations of MUF to produce a range from 0 to 1.2 mM (Ali et al. 2015). Fluorescence was determined ($\lambda_{\text{EX}} = 355 \text{ nm}$, $\lambda_{\text{EM}} = 460 \text{ nm}$) using a Victor3 1420-050 multi-label counter (Perkin Elmer, USA). Enzyme activity was recorded at 25 °C after 0.5, 1, and 2 h (Razavi et al. 2015).

Phenol oxidase activity was analysed following the method described by Saiya-Cork et al. (2002). In brief, fresh soil (0.2 g at 0–10 cm depth) was mixed with 125 mL of 50 mM Tris buffer (pH 7.8) for 2.5 min using a magnetic stirrer. From the resulting suspension, 200 μL was transferred into 96-well microplates, with eight replicate wells per sample for each assay. Subsequently, 50 μL of a 5 mM L-3,4-dihydroxyphenylalanine (L-DOPA) solution was added to each well. The plates were incubated in the dark at 20 °C for 4 h. Absorbance was then measured at 450 nm using a Multi-Mode Microplate Reader (Synergy Mx, BioTek Instruments Inc., USA). Enzyme activity was reported as $\text{mmol h}^{-1} \text{g}^{-1}$ dry soil (Razavi et al. 2015).

2.7 DNA sequencing

Soil DNA was extracted from samples collected at the 0–10 cm layer at the conclusion of the 365-day experiment, using the Zymo Research Soil DNA Kit (Zymo Research, USA). The quantity of extracted DNA was first evaluated through 1% agarose gel electrophoresis, followed by assessment of DNA concentration and purity using a NanoDrop spectrophotometer (Thermo Scientific). Polymerase chain reaction (PCR) amplifications were employed using an ABI 9700 thermocycler (Thermo Fisher Scientific, Waltham, MA, USA). For bacterial community analysis, the V3-V4 region of the 16S rRNA gene (~450–550 bp) was targeted using the primers 341F (CCTAYGGGRBGCASCAG) and 806R (GGACTA CNNGGGTATCTAAT). Fungal community composition was assessed by amplifying the ITS2 region (~380 bp) with primers ITS3 (GCATCGATGAAGAACGCAGC) and ITS4 (TCCTCCGCTTATTGATATGC).

PCR reactions were prepared in a total volume of 25 μL , consisting of 12.5 μL of $2\times$ Taq Plus Master Mix, 3 μL of BSA (2 ng μL^{-1}), 1 μL of each forward and reverse primer (5 μM), 2 μL of template DNA, and 5.5 μL of nuclease-free water. The thermal cycling protocol included an initial denaturation at 95 $^{\circ}\text{C}$ for 5 min, followed by 28 cycles of denaturation at 95 $^{\circ}\text{C}$ for 45 s, annealing at 55 $^{\circ}\text{C}$ for 50 s, and extension at 72 $^{\circ}\text{C}$ for 45 s. A final extension was performed at 72 $^{\circ}\text{C}$ for 10 min, with a subsequent cooling step at 4 $^{\circ}\text{C}$. Amplified products were confirmed through 1% agarose gel electrophoresis and subsequently purified using the Agencourt AMPure XP kit. Paired-end sequencing was performed using the Illumina MiSeq PE300 platform for high-throughput analysis. All PCR amplification and sequencing for bacterial and fungal communities were carried out by Novogene Co., Ltd. (Beijing, China).

2.8 Statistical analyses

Statistical analyses of soil variables across treatments were conducted using both one-way and two-way ANOVA in SPSS version 24 (IBM Corp., Chicago, IL, USA). Before analysis, data were tested for normality and homogeneity of variance using the Shapiro–Wilk and Levene's tests, respectively. When assumptions were met, treatment differences were further examined using Tukey's post-hoc test. For datasets that did not meet these assumptions, \log_{10} or square root transformations were applied as appropriate. Replication ($n=4$) was selected based on constraints typical of mesocosm experiments. A post hoc power analysis indicated that the design had sufficient power ($\geq 80\%$) to detect effect sizes of $d \geq 1.2$ at $\alpha=0.05$. All statistical comparisons were made within individual sampling dates. IBM SPSS version 29 was used to perform ANOVA tests (significance level: $p < 0.05$).

Visualization of temporal trends in soil solution chemistry and other biophysicochemical properties was performed using Origin 2022 (Origin Lab Corp., USA). Microbial community analyses were conducted in R (version 4.4.2; R Core Team, 2024). Microbial community composition differences were assessed using Principal Coordinates Analysis (PCoA) based on Bray–Curtis dissimilarity matrices through the 'vegan' package (Oksanen 2013). Heatmaps representing soil physicochemical parameters and microbial community structure were generated using the 'pheatmap' package to explore relationships between environmental variables and microbial variation.

To assess the influence of specific environmental factors (e.g., pH, NH_4^+ , TN, DOC, total Fe, redox potential, phenol oxidase, β -glucosidase, cellobiase activity, and temperature) on microbial communities, distance-based linear modeling (distLM) was performed using the distLM_forward3 software (Anderson 2003).

Microbial co-occurrence networks were constructed to explore interactions within the soil microbiome. The 300 most abundant bacterial amplicon sequence variants (ASVs) and 200 most abundant Fungal ASVs were selected to minimize spurious correlations while ensuring comparability across treatments. These ASVs represented roughly 54% and 80% of the total bacterial and fungal abundances, respectively. To construct robust microbial co-occurrence networks, pairwise associations between taxa were first calculated using Spearman rank correlation, both of which are suitable for compositional microbial abundance data. To ensure biological relevance and reduce noise, weak or statistically insignificant associations were excluded by applying a correlation threshold ($|r| < 0.6$) and a significance level ($p > 0.05$). The significance of each correlation was further assessed through bootstrap resampling, allowing us to evaluate whether observed interactions deviated from those expected by chance. To account for the large number of comparisons and minimize the risk of false positives, multiple testing correction was performed using the Benjamini–Hochberg false discovery rate (FDR) method. Only statistically significant and strong correlations that passed these criteria were retained for network construction and subsequent topological analysis. Network construction and visualization were performed using Cytoscape version 3.5.1 (Shannon et al. 2003) and Gephi (<https://gephi.github.io/>).

3 Results

3.1 Physicochemical characteristics of soil and pore water

Over the 1-year experiment, soil pH increased in the LW treatment, while biochar addition resulted in a more stable pH of ca. 6.5 during the experimental period

($p < 0.001$; Fig. 1a, b). The Eh of the soil under the low water table regime was relatively stable at +200 mV throughout the experiment (Fig. 1e, f); however, raising the water table slowly decreased the redox potential for the first 30 d after which it remained relatively uniform for the remainder of the experiment (ca. -150 mV) in all treatments, irrespective of Fe or organic C addition (Fig. 2e, f; Fig. S1). Soil solution NO_3^- concentrations were significantly higher in the LW treatments relative to those with a high-water table, especially after the addition of FeSO_4 (Fig. 1g, h). Overall, addition of the organic amendments had no effect on soil NO_3^- concentrations irrespective of FeSO_4 addition. Soil solution DOC concentration was highest in the biochar+HW and C.straw+HW treatments, while the lowest concentrations were observed in the LW treatment (Fig. S1ab). In the HW treatments, NH_4^+ dominated the soluble N pool; however, the concentrations were similar irrespective of amendments and FeSO_4 addition (Fig. S1cd). All the high-water table treatments had a significantly greater soil solution NH_4^+ concentrations relative to the low water table treatment ($p < 0.001$).

The Biochar+HW treatment resulted in the highest aromaticity ($8.05 \text{ L mg C}^{-1} \text{ m}^{-1}$) assessed by specific ultraviolet absorbance at 254 nm (SUVA_{254}) in soil solution, followed by Biochar+HW+ FeSO_4 ($5.2 \text{ L mg C}^{-1} \text{ m}^{-1}$), with the lowest values observed in the Control+LW treatment ($1.3 \text{ L mg C}^{-1} \text{ m}^{-1}$) (Fig. S2b).

3.2 Enzyme activity

The highest β -glucosidase and cellobiase activities were observed in the Paperwaste+HW+ FeSO_4 treatment, followed by the C.straw+HW+ FeSO_4 treatment. In contrast, the Biochar+HW+ FeSO_4 and Biosolid+HW+ FeSO_4 treatments exhibited the most pronounced declines in the activities of both enzymes (Fig. 2a, b). Specifically, β -glucosidase and cellobiase activities in the Biochar+HW+ FeSO_4 were approximately two-fold lower than those in the Control+HW+ FeSO_4 treatment.

Phenol oxidase activity, however, showed a different pattern. The Control+LW+ FeSO_4 recorded the highest activity at $318 \mu\text{mol g}^{-1} \text{ SOM h}^{-1}$, whereas the

Biochar+HW+ FeSO_4 treatment exhibited the lowest value at $124 \mu\text{mol g}^{-1} \text{ SOM h}^{-1}$. Phenol oxidase activity was notably influenced by the water table level, with the (LW condition in the Control+ FeSO_4 treatment significantly enhancing activity. In contrast, phenol oxidase levels remained relatively unchanged under HW treatments, consistently measuring below $200 \mu\text{mol g}^{-1} \text{ SOM h}^{-1}$.

3.3 Extractable soil Fe concentrations

Soil extractable soil Fe(II) concentration dominated over Fe(III) in the HW treatments, while the extractable soil Fe(III) concentrations significantly increased in the Control+LW+ FeSO_4 treatment (Fig. 2f). Correspondingly, the soluble Fe(III) content was 8 g kg^{-1} in the Control+LW+ FeSO_4 , and soil solution Fe(II) concentrations ranged from 8 to 10 g kg^{-1} for all treatments with a high water table at the end of the experiment. Interestingly, Fe bound OC content was greater in Control+LW+ FeSO_4 , whereas there was no obvious change found under the high-water table treatment (Fig. 2d).

3.4 Microbial community characteristics

The relative abundance of dominant microbial taxa shifted in response to soil amendments and water table depth. Under low water table conditions, *Actinobacteria* prevalence was significantly higher, comprising 16–22% of bacterial communities in controls with or without added FeSO_4 (Fig. 3a). Correspondingly, *Ascomycota* dominance reached 50–61% in these treatments, declining in other amendment types (Fig. 3b). Microbial diversity based on the Shannon index spanned 9.4–10.5 for bacteria and 2.5–6.4 for fungi across all treatments (Fig. 3c).

Principal Coordinates Analysis (PCoA) was used to further examine the microbial community structure across treatments, revealing distinct shifts associated with the addition of FeSO_4 (Fig. 3d). The bacterial community composition in the Paperwaste+HW+ FeSO_4 and Control+HW+ FeSO_4 treatments was separated from other treatments, whereas the fungal community in the Paperwaste+HW+ FeSO_4 and C.straw+HW+ FeSO_4 showed the largest separation from Control+HW+ FeSO_4 treatments (Fig. 3d). According to the RDA analysis, pH, EC, cellobiase activity, and redox potential had a strong co

(See figure on next page.)

Fig. 1 Effect of water table manipulation, organic C amendment loaded at 20 t C ha^{-1} and FeSO_4 addition on changes in soil chemical properties in a lowland agricultural peat soil over a one-year in an outdoor mesocosm experiment. Soil pH (a and b), soil EC (c and d), soil redox potential (e and f), nitrate concentration (g and h). The C amendments were loaded at 20 t C ha^{-1} and included biochar (pyrolysed *Miscanthus giganteus* wood chip), commercial paper waste, *M. giganteus* derived chip, barley straw and advanced anaerobically digested biosolids. Water table heights were low water table (LW; -40 cm, representing no change of management), high water table (HW; 0 cm, representing rewetting) and FeSO_4 addition \pm at a rate of 0.5 t ha^{-1} . Values represent mean \pm standard error ($n=4$)

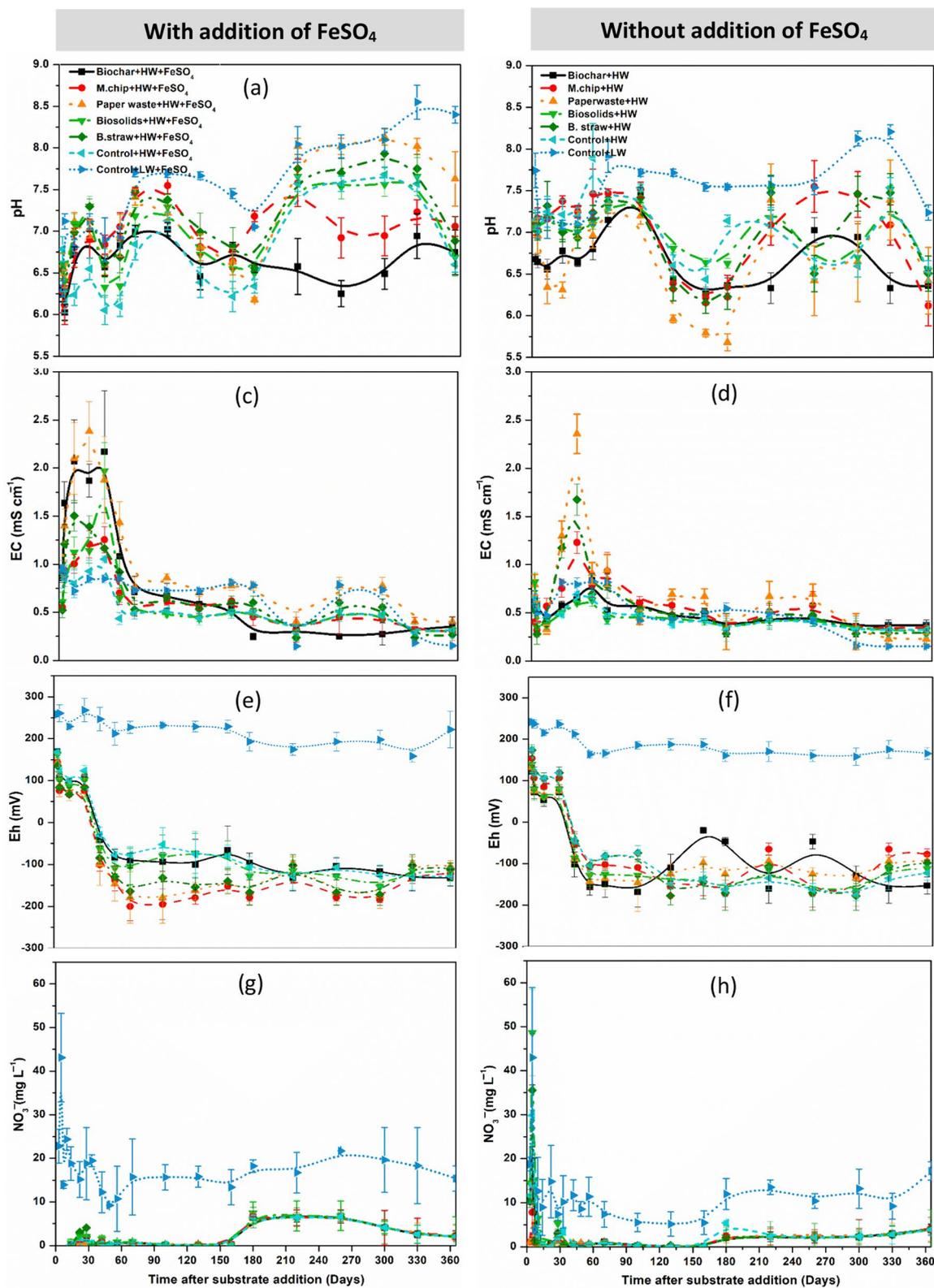


Fig. 1 (See legend on previous page.)

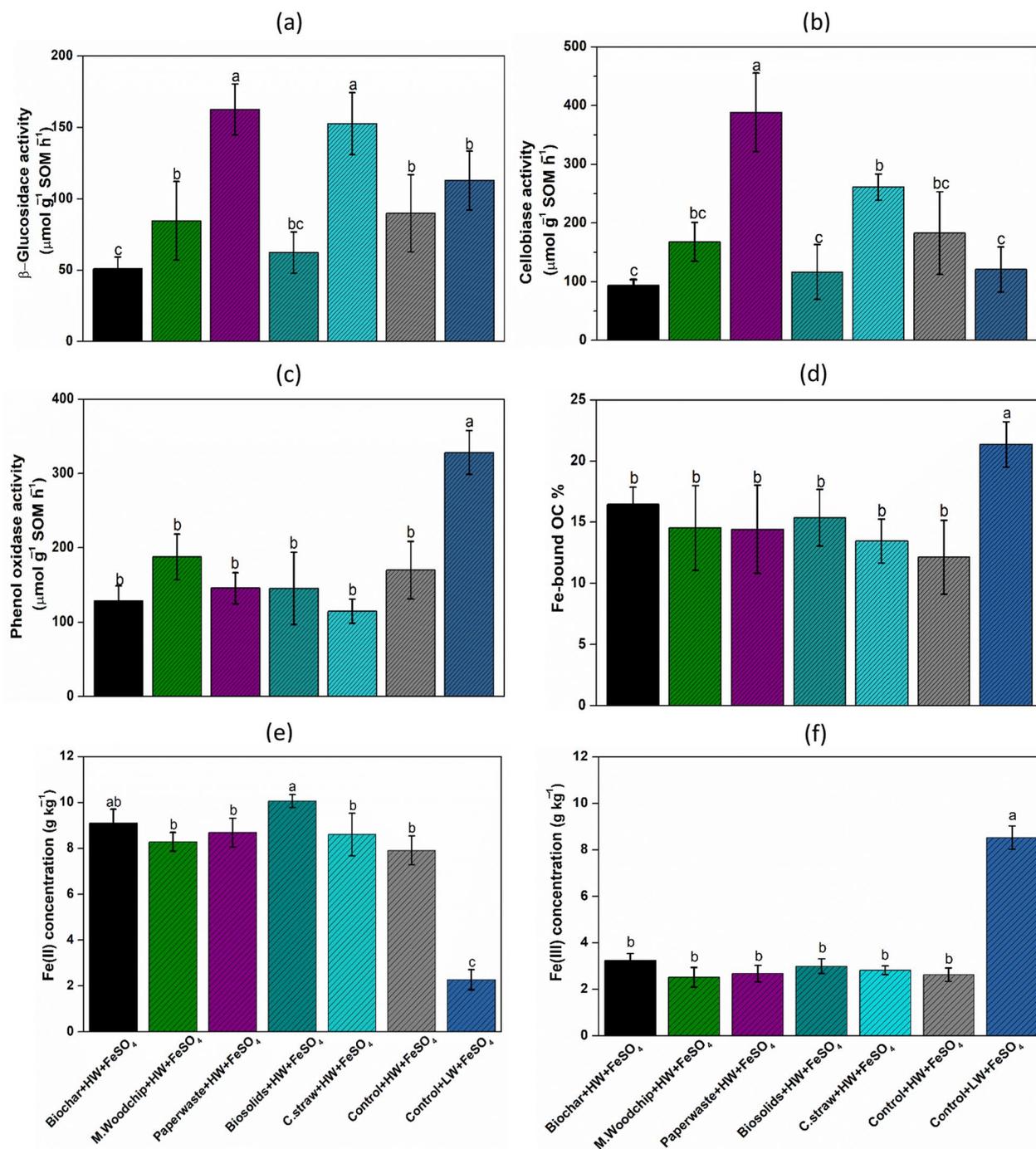


Fig. 2 Effect of organic C amendment and FeSO₄ addition on changes in soil enzyme activities and Fe concentrations in an agricultural peat soil after a year (a) β -glucosidase activity; (b) cellulase activity; (c) phenol oxidative activity; (d) iron-bound soil organic carbon (Fe-bound SOC); (e) Fe(II) concentration; (f) Fe(III) concentration. The C amendments were loaded at 20 t C ha⁻¹ and included biochar (pyrolysed *Miscanthus giganteus* wood chip), commercial paper waste, *M. giganteus* derived chip, barley straw and advanced anaerobically digested biosolids. Water table heights were low water table (LW; -40 cm, representing no change of management), high water table (HW; 0 cm, representing rewetting) and FeSO₄ addition \pm at a rate of 0.5 t ha⁻¹. Values represent mean \pm standard error ($n = 4$)

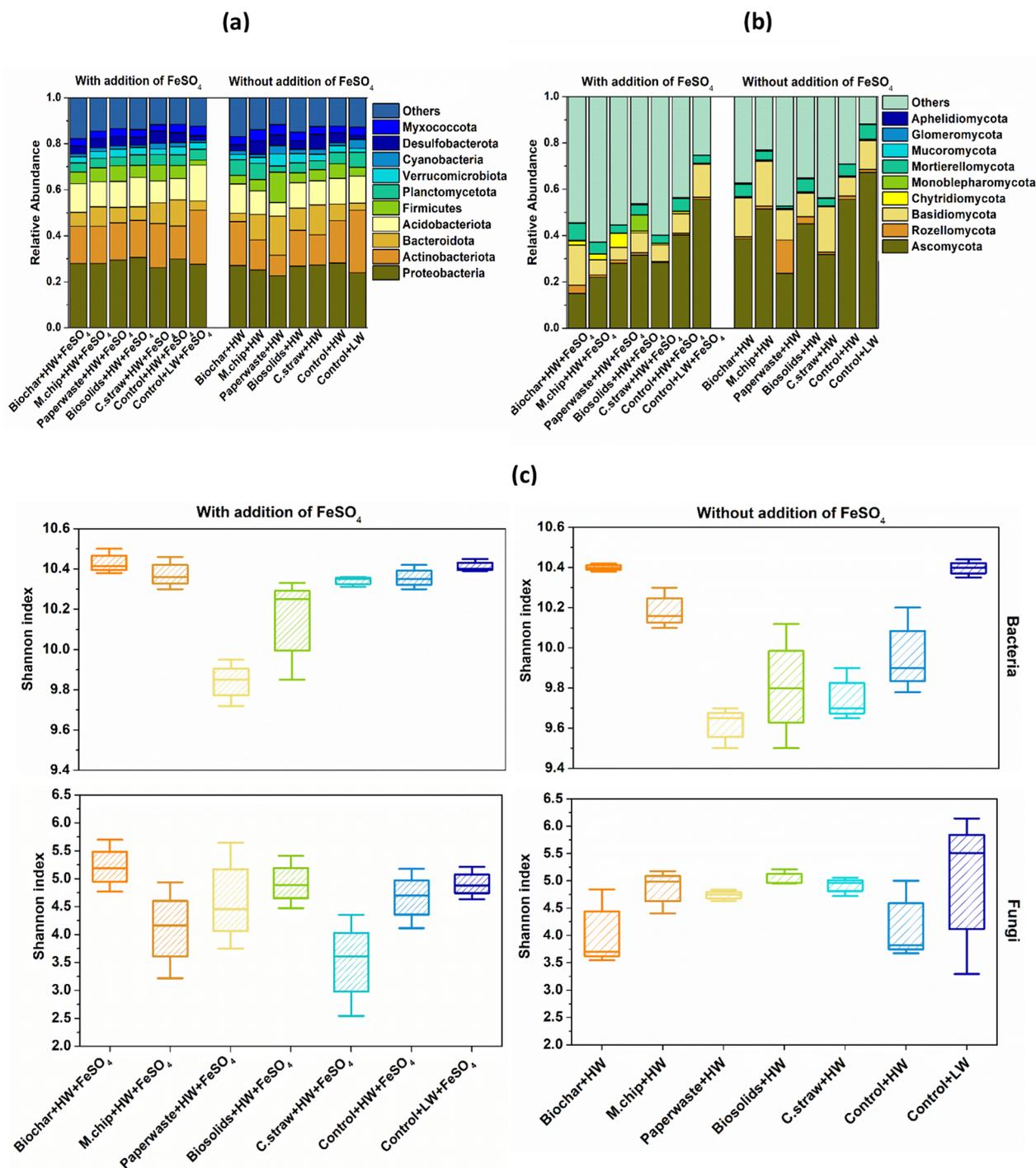


Fig. 3 Bacterial and fungal community dynamics in peat mesocosms. **a** and **b** illustrate the relative abundance of bacterial and fungal community, while **(c)** and **(d)** represent alpha diversity and beta diversity of soil bacterial and Fungal communities, respectively. The Shannon index was calculated with all OTUs. The horizontal bars within boxes represent the median. The tops and bottoms of boxes represent 75th and 25th quartiles, respectively. The principal coordinates analysis (PCoA) with Bray–curtis dissimilarity was performed on the taxonomic profile (at the OTU level) of bacterial communities. The C amendments were loaded at 20 t C ha⁻¹ and included biochar (pyrolysed *Miscanthus giganteus* wood chip), commercial paper waste, *M. giganteus* derived chip, barley straw and advanced anaerobically digested biosolids. Water table heights were low water table (LW; - 40 cm, representing no change of management), high water table (HW; 0 cm, representing rewetting) and FeSO₄ addition ± at a rate of 0.5 t ha⁻¹. Values represent mean ± standard error (n = 4)

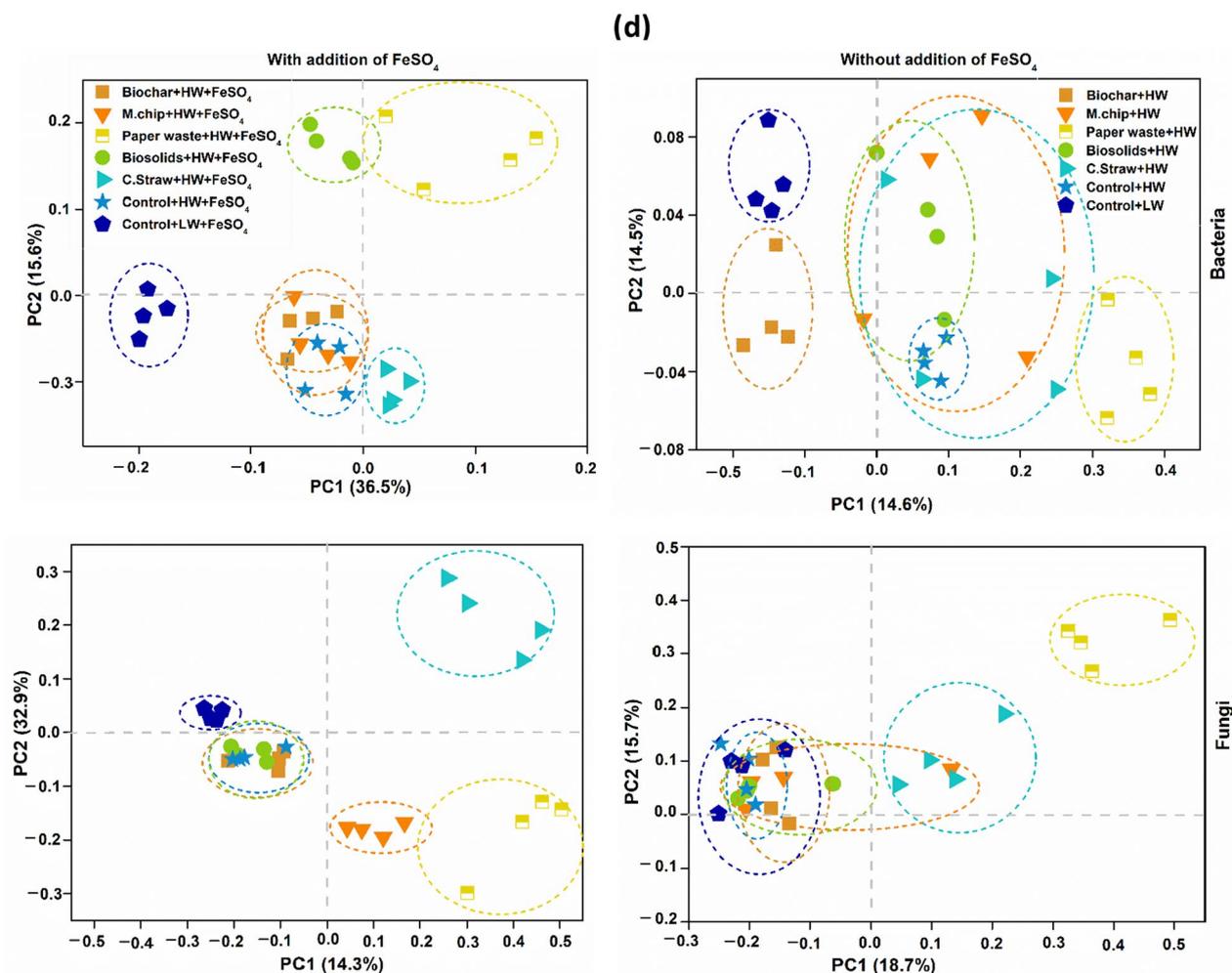


Fig. 3 continued

dependence on both bacterial and fungal community composition (Fig. 5c and d).

Microbial network analysis was performed to discover the interactive relationships between bacterial and fungal communities across treatments. The ratio of positive associations (co-presence) to negative associations (mutual exclusion) was highest in the Control+LW+FeSO₄ treatment and lowest in the Control+HW+FeSO₄ treatment for both microbial groups (Table 1). The fungal network of the Control+LW+FeSO₄ treatment represented 73.1% of nodes affiliated with *Ascomycota*, with the genera *Scutellinia* and *Pyrenochaetopsis* detected as keystone taxa (Table 1). Changes in bacterial and fungal community composition across four treatments were further assessed using Linear Discriminant Analysis Effect Size (LEfSe) (Segata et al. 2011). This approach, applied from the phylum to genus level, enabled the identification of differentially abundant taxa while accounting

for biological consistency and effect size. LEfSe results highlighted specific microbial lineages strongly associated with individual treatments (Fig. 3b), including a notable enrichment of methanogenic genera such as *Methanosarcina* in the Paperwaste + HW treatment.

3.5 Relationships between measured environmental variables and microbial community composition

Distance-based linear model (DistLM) analysis indicated that bacterial community composition was influenced by redox potential (9.2% explanatory power), phenol oxidase activity (6.2%), β-glucosidase activity (7.7%), and total Fe (5.2%). The fungal community was affected by NO₃⁻ concentration (10.6% explanatory power), phenol oxidase activity (8.1%), total Fe (6.5%), NH₄⁺ concentration (5.0%), α-glucosidase activity (4.5%) and β-glucosidase activity (4.7%) (Table 2). Total soil Fe and phenol oxidase activity contributed to shifts in both bacterial and Fungal

Table 1 Properties of bacterial and fungal co-occurrence networks

Properties	Biochar + HW + FeSO ₄		Control + HW + FeSO ₄		Control + LW + FeSO ₄	
	Bacteria	Fungi	Bacteria	Fungi	Bacteria	Fungi
Nodes	170	76	151	69	201	82
Edges	135	69	122	150	164	72
Modularity	0.879	0.835	0.742	0.604	0.807	0.5
ACC	0.121	0.099	0.132	0.200	0.145	0.164
APL	1.181	1.394	1.142	1.644	1.164	1.145
Copresence	82	72	87	84	83	55
Mutual exclusion	18	27	13	16	17	45
Ratio of positive to negative links	4.8	2.6	6.6	5.6	4.5	1.2
Proteobacteria (%)	30.3	34.2	33.7	56.6	30.3	73.1
Actinobacteria (%)	25.8	5.2	18.5	5.8	22	6.1
Acidobacteria (%)	16.9	52.6	17.8	31.8	13	15
Keystone taxa	<i>Pedomicrobium</i> <i>Solirubrobactor</i>	<i>Scutellinia</i>	<i>Anaeromyxobacter</i> <i>ADurb.Bir063-1</i>	<i>Acremonium</i>	<i>Nocardioides</i>	<i>Pyrenochaetopsis</i>

Table 2 Contributions of soil edaphic variables to shaping the bacterial and fungal community based on Bray–Curtis dissimilarities analyzed by distance-based linear modeling (distLM) analysis

Soil edaphic variables	Contribution to bacteria %	Soil edaphic variables	Contribution to fungi %
Redox potential	9.2*	NO ₃ ⁻	10.6***
Phenol oxidase activity	6.2*	Phenol oxidase activity	8.1***
α-glucosidase activity	7.7**	Total Fe	6.5***
Total Fe	5.2*	NH ₄ ⁺	5.0**
β-glucosidase activity	4.5	α-glucosidase activity	4.5*
pH	3.4	β-glucosidase activity	4.7*
NH ₄ ⁺	3.5	Temperature	3.4
Total dissolved N	3.7	pH	3.2
DOC	4.8	Total dissolved N	3.2
Temperature	2.6	DOC	3.3

*, ** and *** indicate a significant result at the $p < 0.05$, $p < 0.01$ and $p < 0.001$ level, respectively

communities, explaining 11.7% and 14.3% of community variation, respectively (Table 2).

Correlation analysis further showed that bacterial community structure was significantly related to phenol oxidase activity, pH, cellobiase activity and DOC (Mantel’s $r \geq 0.2$, $p < 0.01$), while the fungal community correlated with DOC, NO₃⁻, Fe, and enzyme activities

(Mantel’s $r < 0.2$ to 0.4 , $p < 0.01$ to 0.05) (Fig. 5). Co-occurrence networks demonstrated that ratios of positive to negative links between microbes were highest in the Control + LW + FeSO₄ treatment, and fungal nodes were dominated by Ascomycota (73.1% of nodes) for which *Scutellinia* and *Pyrenochaetopsis* were identified as keystone genera (Table 1).

4 Discussion

4.1 Effects of water table management on soil solution chemistry

During the 365-day experimental period, temporal variability of concentration of NO_3^- was significantly higher in the Control+LW+ FeSO_4 treatment compared to all other treatments. The accumulation of NO_3^- observed in this study is likely attributed to aerobic nitrification processes, whereby soil organic matter (SOM) is mineralized, leading to the conversion of organic nitrogen into inorganic forms and the subsequent release of nitrate. Our findings indicate that nitrification predominated under well-aerated conditions, whereas denitrification became the primary nitrogen transformation pathway in treatments characterized by high moisture and anaerobic conditions (Zhu et al. 2013; Huang et al. 2014). These patterns align with previous research on pH-neutral, lowland peat soils, which has shown that lowering the water table promotes nitrification and associated N_2O emissions (Taghizadeh-Toosi et al. 2019), while raising the water table suppresses these processes (Taft et al. 2018). This contrasts with acid upland peat soils where the low soil pH limits can inhibit nitrification (Marsden et al. 2019) were highest among all the treatments, indicating that applying biochar to soil generally leads to an increase in the soil's DOC content, which means that biochar can raise the level of organic C readily available in the soil solution (Nakhavali et al. 2020). This is primarily due to the high C content of biochar itself, and its ability to adsorb and retain organic matter within its porous structure (particularly prevalent in lignocellulosic structures such as *Miscanthus*). After 90 days, the DOC content declined (from 500 to 190 mg L^{-1} soil in FeSO_4 -amended treatment), indicating that leaching of DOC through the mesocosm or sorption of native soil DOM by biochar likely occurred during the experiment period, which might be due to leaching of DOC through the soil profile and the high sorption capacity of biochar for soil DOC (Kasozi et al. 2010; Nakhavali et al. 2020). Additionally, an increase in soil dissolved organic carbon (DOC) content was noted following the application of 500 °C biochar derived from rice straw (Yang et al. 2022a) and wheat straw (Zhang et al. 2017b; Zhang et al. 2021b). These varying outcomes may be attributed to differences in the intrinsic properties of the biochar, which are influenced by feedstock type and pyrolysis conditions. SUVA_{524} values used as a proxy for the aromaticity of DOM are shown in Fig. S2bc, ranging from 5.2 to 8 $\text{L mg C}^{-1} \text{m}^{-1}$. Both biochar treatments, with and without FeSO_4 addition, significantly ($p < 0.05$) elevated SUVA_{524} values compared to the control, indicating enhanced aromaticity of soil DOM (Yang et al. 2021).

During the 365-day experimental period, NO_3^- concentrations exhibited significantly greater temporal variability in the Control+LW+ FeSO_4 treatment compared to all other treatments. This pronounced NO_3^- accumulation is best explained by aerobic nitrification of SOM, where microbial oxidation of organic nitrogen results in the release of NO_3^- . Under low moisture and oxic conditions, nitrification appeared to be the dominant nitrogen transformation pathway. In contrast, treatments with higher water tables created more anaerobic conditions, which favored denitrification as the primary nitrogen pathway (Zhu et al. 2013; Huang et al. 2014; Jeewani et al. 2025). These findings are supported by prior studies in lowland, pH-neutral peat soils, where lowering the water table increased N_2O emissions via enhanced nitrification, while raising the water table suppressed these emissions (Taghizadeh-Toosi et al. 2019; Taft et al. 2018). In contrast, nitrification is less prominent in acid upland peat soils due to pH limitations (Marsden et al. 2019).

These results support the hypothesis that nutrient stoichiometry specifically the balance between carbon and nitrogen availability modulates microbial activity and decomposition processes in peat soils. Under saturated (anaerobic) conditions, microbial access to oxygen is restricted, and nitrogen becomes a co-limiting factor, particularly when high C:N ratio substrates (e.g., Biochar+HW) are present. In contrast, under aerobic conditions and with lower C:N ratios, nitrogen availability increases, stimulating nitrifier populations and associated enzyme activities. Thus, the data reveal a mechanistic link between hydrological conditions, nutrient availability, and microbial nitrogen cycling, validating the hypothesis that stoichiometric constraints regulate microbial functioning in peatland environments.

Moreover, DOC concentrations were highest in the biochar-amended treatments (with and without FeSO_4), highlighting biochar's capacity to elevate dissolved organic carbon in soil solutions. This increase can be attributed to biochar's high C content and its porous structure, which allows for adsorption and retention of both added and native organic matter (Nakhavali et al. 2020). However, DOC levels declined after 90 days (e.g., from 500 to 190 mg L^{-1} in Fe-amended treatments), likely due to leaching losses and sorption of native dissolved organic matter (DOM) by biochar (Kasozi et al. 2010). The observed differences across studies (e.g., Yang et al. 2022b; Zhang, et al., 2017a); Zhang, et al. 2021a) may reflect variations in feedstock type and pyrolysis conditions, which influence the physicochemical properties of biochar such as total organic carbon, aromaticity, and porosity pore connectivity (Keiluweit et al. 2010; Kloss et al. 2012).

Additionally, both biochar treatments significantly increased SUVA₂₅₄ values (ranging from 5.2 to 8 L mg C⁻¹ m⁻¹), an indicator of DOM aromaticity (Fig. S2bc), suggesting that biochar enhanced the aromatic character of soil DOC (Yang et al. 2022b). These changes further support the view that biochar influences not only carbon quantity but also carbon quality in peat systems, with potential implications for microbial processing and carbon stability.

4.2 Microbial responses to a change in water levels and organic amendments

Bacteria and fungi play distinct but complementary roles in peat carbon decomposition (Kluber et al. 2020). In our study, variations in microbial abundance and composition were strongly driven by changes in water table levels and amendment quality, emphasizing the influence of C resource traits on microbial ecological strategies and peat decomposition. Microbial abundance patterns confirmed that LW conditions favored aerobic metabolism, as reflected in significantly increased bacterial and fungal populations (Fig. 3a, b). The highest abundance of *Actinobacteria*, observed in the Control+LW treatments regardless of FeSO₄ addition (Fig. 2), suggests that dry, oxic environments with relatively labile organic matter support r-strategist bacterial taxa capable of rapid resource exploitation (Richy et al. 2024). These bacteria are also known to stimulate SOC mineralization via microbial priming, due to their mycelial morphology and ability to access otherwise protected C pools (Luo et al. 2017; Fu et al. 2022). This coincided with the highest recorded CO₂ emissions (17.4 t CO₂e ha⁻¹ yr⁻¹; Jeewani et al. 2025), likely the result of enhanced microbial turnover and accelerated carbon cycling (Luo et al. 2011).

Fungal community shifts further supported the resource-quality-driven life history response. Ascomycota, a phylum composed largely of K-strategists with high substrate versatility and stress tolerance, dominated the Control+LW treatment (68% relative abundance). These fungi were significantly associated with oxidative enzyme activity especially phenol oxidase (8.1; $p < 0.01$; Fig. 4, Table 2) indicating their role in degrading aromatic, complex organic matter. The microbial network in this treatment was also dominated by Ascomycota (73.1%), including cellulolytic keystone genera such as *Scutellinia* and *Pyrenochaetopsis*, known to contribute to the mineralization of stable C compounds (Zhang et al. 2021a). This dominance aligns with increased CO₂ flux and further confirms that K-strategist fungi are selectively enriched in conditions where C availability is high but structurally complex (Yi et al. 2019; Boer et al. 2005).

In contrast, saturated, HW conditions especially in the Paper waste+HW treatment favored anaerobic C

pathways. CH₄ emissions were significantly elevated (Table 1; Jeewani et al. 2025), coinciding with increased abundance of the acetoclastic methanogen *Methanosarcina* (Fig. 4b). This suggests that under reducing conditions, r-strategist methanogens exploit labile C inputs from cellulose-rich amendments. Supporting studies have shown that acetate-using (*Methanotrichaceae*) and hydrogenotrophic (*Methanobacteriaceae*) methanogens dominate methane production in organic-rich peat systems (Breeuwer et al. 2009; Corbett et al. 2015; Bräuer et al. 2020).

Biochar application had a marked effect on microbial life strategies and ecosystem-level C and N cycling. In biochar-amended soils, the relative abundance of Ascomycota declined significantly, and microbial co-occurrence networks were simplified (Figs. 3a–d, 4a), indicating that biochar altered the ecological niches available to K-strategist fungi. This shift likely reflects changes in DOC quality particularly reductions in labile C and increases in aromatic C resulting from biochar's sorptive capacity and structural properties (Zimmerman 2010; Han et al. 2016). The “charosphere” environment also promoted microbial carbon use efficiency and moderated community composition (Quilliam et al. 2013; Lehmann et al. 2011), favoring generalist taxa and possibly enhancing nutrient immobilization (Fig. 5).

Functionally, biochar significantly suppressed CH₄ emissions by 20.7 t CO₂e ha⁻¹ yr⁻¹ (Jeewani et al. 2025), likely through inhibition of methanogenic activity. This suppression was linked to a lower ratio of methanogen (*mcrA* gene) to methanotroph (*pmoA* gene) abundance (Dong et al. 2013; Han et al. 2016), a pattern consistent with reduced availability of easily fermentable substrates. Thus, biochar appears to act as a selective filter, modifying resource quality and shaping microbial community assembly and function in ways that suppress r-strategist anaerobes and alter K-strategist fungal dominance.

Furthermore, nitrifying microbial genera (e.g., *Nitrospira*, *Nitrosomonas*, *Nitrosovibrio tenuis*, *Bacillus* spp.) were dominant under low water table conditions (Fig. 4b, Table 2), corresponding with high NO₃⁻ availability and N₂O emissions (Jeewani et al. 2025). These taxa, adapted to oxic, nutrient-rich environments, exemplify r-strategist behavior by rapidly exploiting available NH₄⁺ and promoting nitrification-driven N losses. The DistLM model corroborated these findings, showing strong associations between microbial community shifts and environmental variables such as redox potential (9.2; $p < 0.01$) and NO₃⁻ content (10.6; $p < 0.001$).

Collectively, these results confirm that resource quality not just its quantity plays a pivotal role in structuring microbial life strategies in peat soils. Under nutrient-rich and oxic conditions, r-strategists dominate and promote

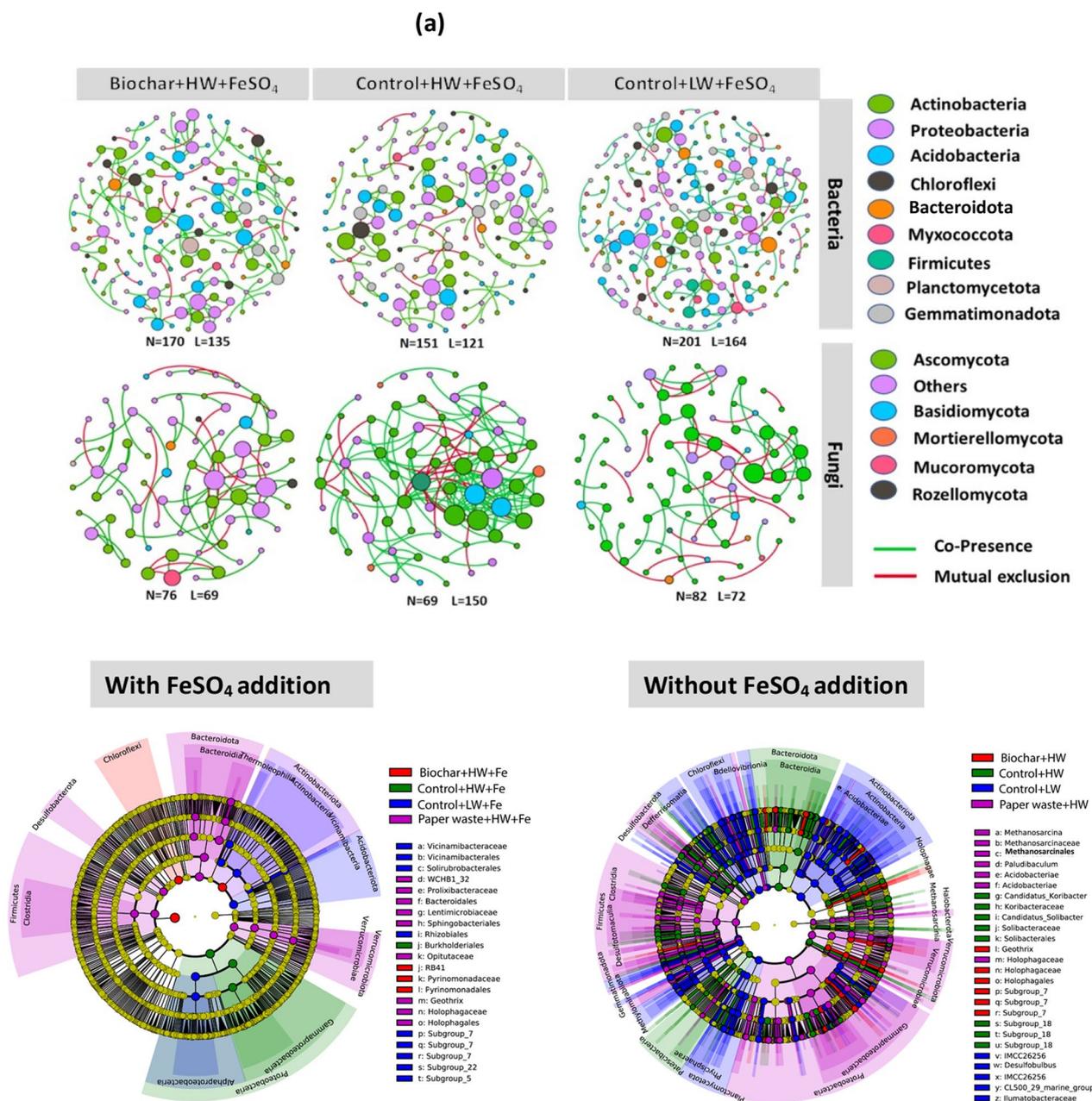


Fig. 4 Co-occurrence networks of the abundance OTU (a) N, node. L, line. Edges represent significant Spearman correlations ($\rho > |0.8|, p < 0.05$). Red Lines represent a significant negative correlation, and green Lines represent a significant positive. Changes in bacterial community composition between the selected 4 treatments out of all treatments (Linear discriminant analysis Effect Size) explain differences between classes by coupling standard tests for statistical significance with additional tests encoding biological consistency and effect relevance (b). The C amendments were loaded at 20 t C ha⁻¹ and included biochar (pyrolysed *Miscanthus giganteus* wood chip), commercial paper waste, *M. giganteus* derived chip, barley straw and advanced anaerobically digested biosolids. Water table heights were low water table (LW; - 40 cm, representing no change of management), high water table (HW; 0 cm, representing rewetting) and FeSO₄ addition ± at a rate of 0.5 t ha⁻¹. Values represent mean ± standard error (n=4)

rapid mineralization, while biochar moderates these responses by constraining labile resource availability and selectively shaping community composition. These insights offer a mechanistic basis for understanding how

peatland management strategies such as biochar addition can modulate decomposition pathways and greenhouse gas emissions by altering microbial resource use strategies.

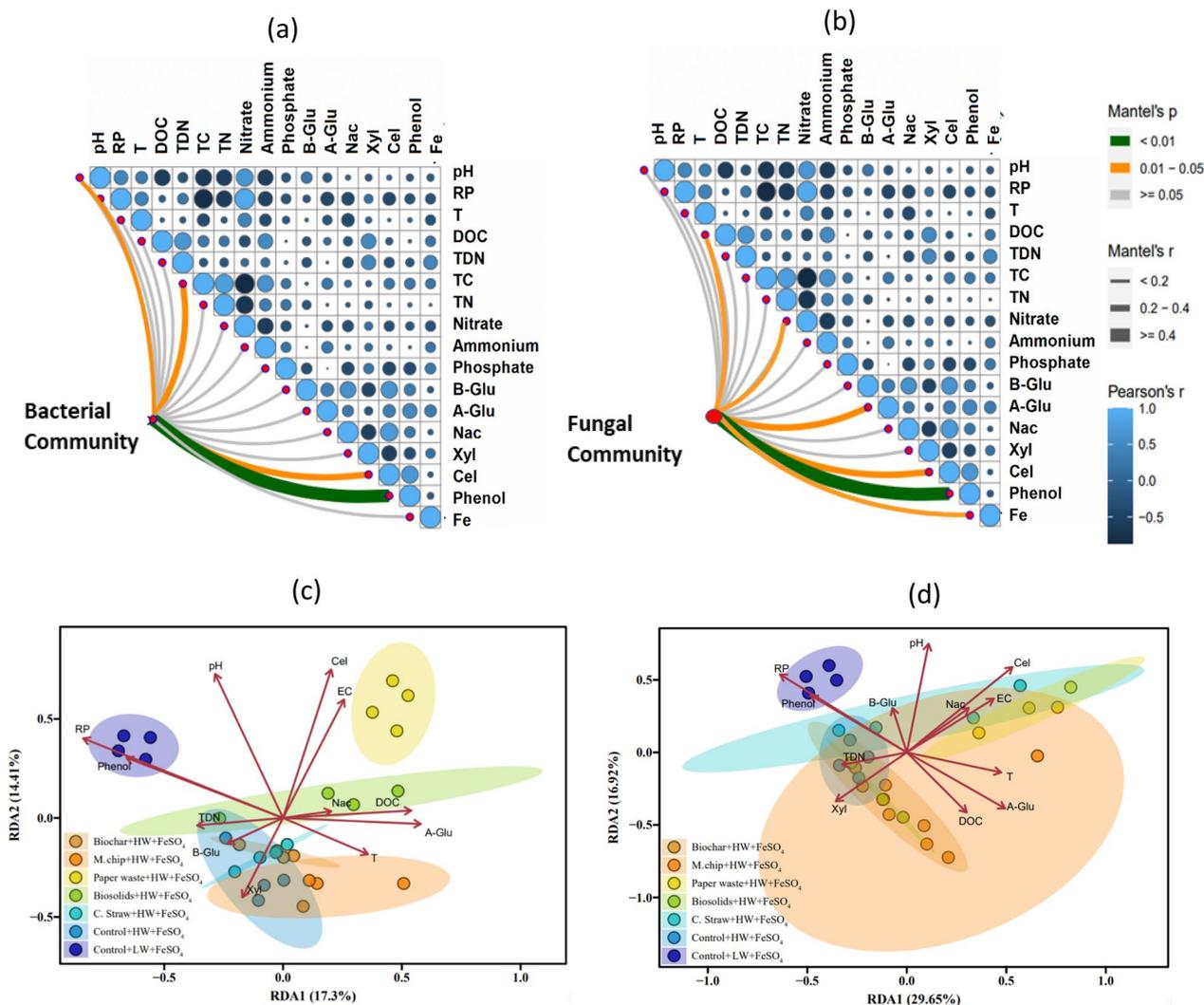


Fig. 5 Correlations of the bacterial and fungal community composition (Bray–Curtis distance) with soil variables (a and b, respectively). Edge width corresponds to the Mantel's r value, and the edge colour denotes the statistical significance of these variables are shown with a colour gradient denoting Pearson's correlation coefficient. c and d represent redundancy analysis (RDA) showing the relationship between the soil bacterial (e) and fungal (f) community and environmental factors. Abbreviation for soil variables include pH, RP, redox potential; T, Temperature; DOC, dissolved organic carbon; TDN, total dissolved nitrogen; TC, total carbon; TN, total nitrogen; Nitrate; NO_3^- , Ammonium; NH_4^+ , phosphate; PO_4^{3-} , B-Glu; β -glucosidase, Nac; N-acetyl- β lucosaminidases, Xyl; β -Xylosidases, Cel; β -Cellulobiosidase, Phenol; phenol oxidase and Fe; total Fe concentration. The C amendments were loaded at 20 t C ha^{-1} and included biochar (pyrolysed *Miscanthus giganteus* wood chip), commercial paper waste, *M. giganteus* derived chip, barley straw and advanced anaerobically digested biosolids. Water table heights were low water table (LW; - 40 cm, representing no change of management), high water table (HW; 0 cm, representing rewetting) and FeSO_4 addition \pm at a rate of 0.5 t ha^{-1} . Values represent mean \pm standard error ($n = 4$)

4.3 Trade-offs between the 'enzyme latch' and 'iron gate' mechanisms

Soil enzymes play a main role in SOM decomposition (Soares and Rousk 2019). The soil enzyme activities among organic amendments, as well as water table levels, differed significantly (Fig. 2a–c), indicating that they were significantly affected by availability of organic C and hydrological management practices

(Fig. S4). Hydrolytic enzyme activities (β -glucosidase and cellobiohydrolase) were significantly higher in the Paper Waste+HW+ FeSO_4 and C.straw treatments compared to the Control+HW+ FeSO_4 treatment. This enhancement may be attributed to the substantial increase in available carbon substrates for microbial mineralization. Previous studies also reported lower soil enzyme activities in degraded agricultural peatland

with reduced C source than those in natural peatland, thus limiting soil enzyme activity (Wang et al. 2021).

However, phenol oxidative activity increased in the low water table peat mesocosms upon exposure to O₂ (Freeman et al. 1996, 2001b). In contrast, several studies have observed that the presence of Fe(II) in hypoxic peatland soils may enhance phenol oxidative activity (Li et al. 2012; Hall and Silver 2013). However, the activity of polyphenol oxidase and Fe(III) content in Control + LW + FeSO₄ was higher than in other treatments (Fig. 2c). This may be due to the low soil water content and good soil air permeability in the soil core, which are favorable to the oxidation of Fe(II) and the survival and reproduction of aerobic microorganisms, thereby promoting degradation of phenolic compounds and facilitating SOM decomposition (Freeman et al. 1996, 2004; Romanowicz et al. 2015). This finding indicates that oxygen availability is the primary factor influencing phenol oxidase activity in this study.

In contrast, HW treatments resulted in reducing conditions with low phenol oxidase activity. These results are consistent with the mechanism associated with the “enzyme latch” from previous studies (Freeman et al. 2001a; Fenner and Freeman 2011; Wang et al. 2024b).

Fe(II) oxidation plays a key role in biochemical processes during water table fluctuation in peat soils that involves the mobilization and stabilization of C (Li et al. 2012; Wang et al. 2024b). It is noteworthy that Fe(III) concentration and Fe bound OC were greater in the Control + LW + FeSO₄ treatment cores ($p > 0.05$; Fig. 2d), which is consistent with previous publications regarding the “iron gate” mechanism (Wang et al. 2017). In drained peat cores, soluble Fe(II) is oxidized to less soluble Fe(III), enhancing SOM preservation. Furthermore, SOM may absorb or co-precipitate with the newly-formed reactive Fe(III) oxides, forming Fe-bound OC complexes that are chemically more stable (Kaiser and Guggenberger 2007; Riedel et al. 2013; Feng et al. 2025). Such associations between iron and carbon are thought to play a key role in stabilizing and preserving soil C (Kaiser and Guggenberger 2007; Lalonde et al. 2012). In particular, aromatic and phenolic compounds that accumulate in wetland environments have been shown to bind strongly to iron oxides (Lalonde et al. 2012; Feng et al. 2025), potentially protecting them from decomposition in the presence of O₂ (Hall et al. 2016). Full quantification of C balance in soil cores with low and high-water table levels offers the unique opportunity to explore the trade-offs between the “enzyme latch” and “iron gate” mechanisms.

4.4 Implications

Sustainable management strategies that increase C stocks while simultaneously reducing GHG emissions are needed to restore drained peatlands. Large-scale

rewetting is currently regarded as the most effective strategy for halting further degradation and subsidence of peat soils, while also mitigating GHG emissions—particularly CO₂ (Evans et al. 2021). While rewetting is a critical step in the ecological recovery of drained peatlands, its success in reducing overall GHG emissions is strongly influenced by the interplay between soil geochemistry and resident microbial communities. Sustaining a high-water table in agricultural peatlands tends to enhance the abundance and potentially the activity of methanogenic microbes, which can in turn result in increased CH₄ emissions (Jeewani et al. 2025). The presence of terminal electron acceptors such as Fe(III) and SO₄²⁻ inhibits methanogenesis. Oxic conditions promote nitrification by enhancing the activity of nitrifying bacteria, which can result in increased short-term N₂O emissions depending on the geochemical characteristics and microbial community structure (Wang et al. 2024a). The modified microbial abundance, diversity and co-occurrence pattern towards reduced abundance of C decomposers in the Biochar + HW + FeSO₄ treatment suggests that this combination of treatments shows significant potential for increasing the C stabilised in the peat, through the suppression of CH₄ emissions in addition to adding recalcitrant C (Fig. 6). As such it is highly recommended that such amendments should be trialed at a larger field scale over longer timeframes.

5 Conclusion

This study investigated the biogeochemical dynamics of agricultural peat soils under various organic and inorganic amendment strategies alongside with water table management. Our results demonstrate that maintaining a high-water table combined with biochar is a critical factor in suppressing key decomposer microbial taxa, particularly Ascomycota and Actinobacteria, along with their associated enzymatic activities. Amendments with low C:N ratios such as biosolids, straw and paper waste significantly promoted r-strategist microbial populations by alleviating limitations in labile carbon and nitrogen availability and altering microbial community composition. Notably, the combination of an elevated water table and paper waste enhanced methanogenic activity, while the application of high C:N biochar under the same conditions suppressed Ascomycota abundance. Moreover, the integrated application of biochar, an elevated water table, and FeSO₄ (Biochar + HW + FeSO₄) substantially restructured fungal and bacterial co-occurrence networks via modified biogeochemistry of peat soil. Collectively, these findings suggest that the Biochar + HW + FeSO₄ treatment regime effectively alters microbial interactions and biogeochemical processes, contributing to reduced

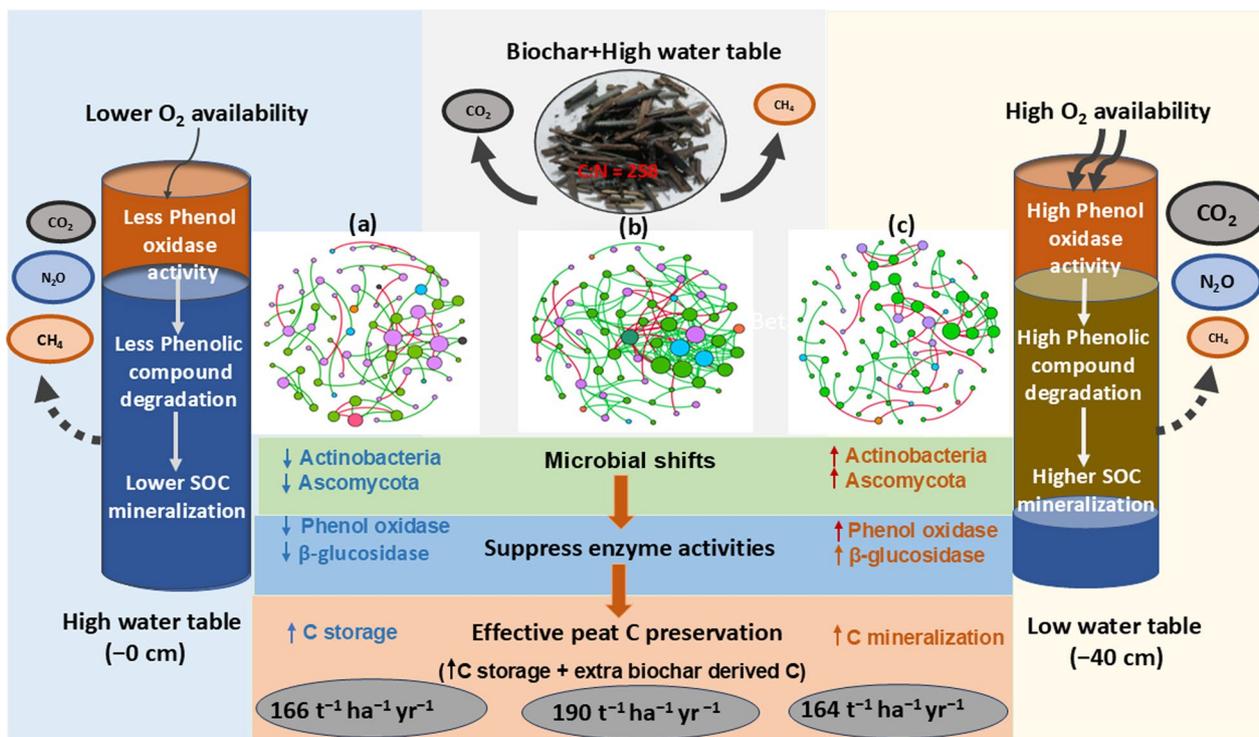


Fig. 6 Proposed mechanisms for microbially mediated of soil organic carbon (SOC) preservation in peat soil alongside with rewetting and biochar amendments after 360 days of mesocosm experiment. Significant difference in C pools between three main treatments such as Control + high water table, Control + low water table and Biochar + high water table $p < 0.05$. **a, b and c** indicate fungal co-occurrence networks of each treatment and its complexity (nodes represent different phylum and lines represent negative (green) and positive (red) interactions). Numbers indicate new C retention in each pool based on Jeewani et al. (2025)

organic matter decomposition and GHG emissions. This integrated management approach offers a promising strategy for enhancing the sustainability and climate resilience of agricultural peatlands.

Supplementary Information

The online version contains supplementary material available at <https://doi.org/10.1007/s42773-025-00501-y>.

Additional file 1.

Acknowledgements

We would like to thank James Brown and John Taylor for hosting and providing access to the field experiments from which samples were collected, and for their wider support to the project. Further, we thank Joe Cotton assistance with sample collection in addition to Dŵr Cymru Welsh Water and Ahlstrom for providing the biosolids and paper waste, respectively.

Author contributions

JHP: Fieldwork, Investigation, Methodology, Formal analysis, Visualization, Writing—Original Draft; RWB: Writing—Review & Editing; JR: Review & Editing, CDE: Conceptualization, Writing—Review & Editing, Funding acquisition; DRC: Conceptualization, Supervision, Writing—Review & Editing, Funding acquisition; DLJ: Conceptualization, Supervision, Writing—Review & Editing, Funding acquisition. All authors read and approved the final manuscript.

Funding

This work was supported by UK Research and Innovation (UKRI) through the Biotechnology and Biological Sciences Research Council (BBSRC) funded Greenhouse Gas Removal Peatland Demonstrator project (BB/V011561/1).

Data availability

Data available within the article or its supplementary materials.

Declarations

Competing interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Author details

¹School of Environmental and Natural Sciences, Bangor University, Bangor, Gwynedd LL57 2UW, UK. ²UK Centre for Ecology and Hydrology, Bangor, Gwynedd LL57 2UW, UK.

Received: 9 May 2025 Revised: 21 July 2025 Accepted: 24 July 2025

Published online: 10 September 2025

References

Ali RS, Ingwersen J, Demyan MS, Funkuin YN, Wizemann H-D, Kandeler E, Poll C (2015) Modelling in situ activities of enzymes as a tool to explain seasonal variation of soil respiration from agro-ecosystems. *Soil Biology and Biochemistry* 81:291–303

- Andersen R, Chapman SJ, Artz RRE (2013) Microbial communities in natural and disturbed peatlands: a review. *Soil Biol Biochem* 57:979–994
- Anderson MJ (2003) DISTLM forward: a FORTRAN computer program to calculate a distance-based multivariate analysis for a linear model using forward selection. University of Auckland, New Zealand, Department of Statistics, p 10
- Ascough PL, Bird MI, Brock F, Higham TFG, Meredith W, Snape CE, Vane CH (2009) Hydroxypropylation as a new tool for radiocarbon pre-treatment and the quantification of black carbon. *Quat Geochronol* 4:140–147
- Blake, G.R., Hartge, K.H., 1986. Particle Density, *Methods of Soil Analysis*, pp. 377–382
- Bonn A, Allott T, Evans M, Joosten H, Stoneman R (2016) Peatland restoration and ecosystem services: science, policy and practice. Cambridge University Press
- Bradfield EG, Cooke DT (1985) Determination of inorganic anions in water extracts of plants and soils by ion chromatography. *Analyst* 110:1409–1410
- Bräuer LS, Basiliko N, MP Siljanen HH, Zinder S (2020) Methanogenic archaea in peatlands. *FEMS Microbiology Letters* 367, fnaa172
- Breeuwer A, Robroek BJM, Limpens J, Heijmans MMPD, Schouten MGC, Berendse F (2009) Decreased summer water table depth affects peatland vegetation. *Basic and Applied Ecology* 10:330–339
- Brookes PC, Landman A, Pruden G, Jenkinson DS (1985) Chloroform fumigation and the release of soil nitrogen: a rapid direct extraction method to measure microbial biomass nitrogen in soil. *Soil Biology and Biochemistry* 17:837–842
- Brouns K, Keuskamp JA, Potkamp G, Verhoeven JTA, Hefting MM (2016) Peat origin and land use effects on microbial activity, respiration dynamics and exo-enzyme activities in drained peat soils in the Netherlands. *Soil Biol Biochem* 95:144–155
- Brown RW, Chadwick DR, Bending GD, Collins CD, Whelton HL, Daulton E, Covington JA, Bull ID, Jones DL (2022) Nutrient (C, N and P) enrichment induces significant changes in the soil metabolite profile and microbial carbon partitioning. *Soil Biol Biochem* 172:108779
- Butterbach-Bahl K, Baggs EM, Dannenmann M, Kiese R, Zechmeister-Boltenstern S (2013) Nitrous oxide emissions from soils: how well do we understand the processes and their controls? *Philos Trans Royal Soc B Biol Sci* 368:20130122
- Chowdhury, Z.K., 2013. Activated carbon: solutions for improving water quality. American Water Works Association
- Corbett JE, Tfaily MM, Burdige DJ, Glaser PH, Chanton JP (2015) The relative importance of methanogenesis in the decomposition of organic matter in northern peatlands. *Journal of Geophysical Research: Biogeosciences* 120:280–293
- Dean JF, Middelburg JJ, Röckmann T, Aerts R, Blauw LG, Egger M, Jetten MSM, de Jong AEE, Meisel OH, Rasigraf O (2018) Methane feedbacks to the global climate system in a warmer world. *Rev Geophys* 56:207–250
- Deshoux M, Sadet-Bourgeteau S, Gentil S, Prévost-Bouré NC (2023) Effects of biochar on soil microbial communities: a meta-analysis. *Sci Total Environ* 902:166079
- Dong D, Yang M, Wang C, Wang H, Li Y, Luo J, Wu W (2013) Responses of methane emissions and rice yield to applications of biochar and straw in a paddy field. *Journal of Soils and Sediments* 13:1450–1460
- Wang Y, Knorr KH, Sun J, Sun D, Xu G, Li H, Wu F, Xu Z, Wang S (2024b) Effect of water table restoration on microbial communities and enzyme activities in drained peatland. *Land Degradation Dev*
- Eilers KG, Lauber CL, Knight R, Fierer N (2010) Shifts in bacterial community structure associated with inputs of low molecular weight carbon compounds to soil. *Soil Biol Biochem* 42:896–903
- Evans CD, Peacock M, Baird AJ, Artz RRE, Burden A, Callaghan N, Chapman PJ, Cooper HM, Coyle M, Craig E (2021) Overriding water table control on managed peatland greenhouse gas emissions. *Nature* 593:548–552
- Evans C, Artz R, Moxley J, Smyth M-A, Taylor E, Archer E, Burden A, Williamson J, Donnelly D, Thomson A (2017) Implementation of an emissions inventory for UK peatlands. Centre for Ecology and Hydrology
- Evans CD, Morrison R, Cumming A, Bodo A, Burden A, Callaghan N, Clilverd H, Cooper H, Cowan N, Crabtree D (2023) Defra Lowland Peat 2: Managing agricultural systems on lowland peat for decreased greenhouse gas emissions whilst maintaining agricultural productivity. Report to Defra for Project SP1218
- Feng X, Zhao Y, Wang H, Liu C (2025) Iron-organic carbon interactions in wetlands: implications for wetland carbon preservation under global changes. *Glob Change Biol* 31:e70300
- Fenner N, Freeman C (2011) Drought-induced carbon loss in peatlands. *Nat Geosci* 4:895–900
- Freeman C, Liska G, Ostle NJ, Lock MA, Reynolds B, Hudson J (1996) Microbial activity and enzymic decomposition processes following peatland water table drawdown. *Plant Soil* 180:121–127
- Freeman C, Ostle N, Kang H (2001a) An enzymic “latch” on a global carbon store. *Nature* 409:149–149
- Freeman C, Ostle NJ, Fenner N, Kang H (2004) A regulatory role for phenol oxidase during decomposition in peatlands. *Soil Biol Biochem* 36:1663–1667
- Fu Y, Luo Y, Auwal M, Singh BP, Van Zwieten L, Xu J (2022) Biochar accelerates soil organic carbon mineralization via rhizodeposit-activated Actinobacteria. *Biology and Fertility of Soils* 58:565–577
- Hall SJ, Silver WL (2013) Iron oxidation stimulates organic matter decomposition in humid tropical forest soils. *Glob Change Biol* 19:2804–2813
- Hall SJ, Silver WL, Timokhin VI, Hammel KE (2016) Iron addition to soil specifically stabilized lignin. *Soil Biol Biochem* 98:95–98
- Han X, Sun X, Wang C, Wu M, Dong D, Zhong T, Thies JE, Wu W (2016) Mitigating methane emission from paddy soil with rice-straw biochar amendment under projected climate change. *Scientific Reports* 6:24731
- Hu J, Wu H, Sun Z, Peng Q-A, Zhao J, Hu R (2020) Ferrous iron addition decreases methane emissions induced by rice straw in flooded paddy soils. *ACS Earth Space Chem* 4:843–853
- Huang Y, Li Y, Yao H (2014) Nitrate enhances N₂O emission more than ammonium in a highly acidic soil. *Journal of Soils and Sediments* 14:146–154
- Jeewani PH, Brown RW, Rhymes JM, McNamara NP, Chadwick DR, Jones DL, Evans CD (2025) Greenhouse gas removal in agricultural peatland via raised water levels and soil amendment. *Biochar* 7:39
- Jeewani PH, Brown RW, Evans CD, Cook J, Roberts BP, Fraser MD, Chadwick DR, Jones DL (2025) Rewetting alongside biochar and sulphate addition mitigates greenhouse gas emissions and retain carbon in degraded upland peatlands. *Soil Biology and Biochemistry* 207
- Jones DL, Rousk J, Edwards-Jones G, DeLuca TH, Murphy DV (2012) Biochar-mediated changes in soil quality and plant growth in a three year field trial. *Soil Biology and Biochemistry* 45:113–124
- Kaiser K, Guggenberger G (2007) Sorptive stabilization of organic matter by microporous goethite: sorption into small pores vs. surface complexation. *Eur J Soil Sci* 58:45–59
- Kandel TP, Lærke PE, Hoffmann CC, Elsgaard L (2019) Complete annual CO₂, CH₄, and N₂O balance of a temperate riparian wetland 12 years after rewetting. *Ecol Eng* 127:527–535
- Kasazi GN, Zimmerman AR, Nkedi-Kizza P, Gao B (2010) Catechol and humic acid sorption onto a range of laboratory-produced black carbons (biochars). *Environ Sci Technol* 44:6189–6195
- Keilulweit M, Nico PS, Johnson MG, Kleber M (2010) Dynamic molecular structure of plant biomass-derived black carbon (Biochar). *Environ Sci Technol* 44:1247–1253
- Kloss S, Zehetner F, Dellantonio A, Hamid R, Ottner F, Liedtke V, Schwanninger M, Gerzabek MH, Soja G (2012) Characterization of slow pyrolysis biochars: effects of feedstocks and pyrolysis temperature on biochar properties. *J Environ Qual* 41:990–1000
- Kluber LA, Johnston ER, Allen SA, Hendershot JN, Hanson PJ, Schadt CW (2020) Constraints on microbial communities, decomposition and methane production in deep peat deposits. *PLoS ONE* 15:e0223744
- Koch O, Tschirko D, Kandeler E (2007) Temperature sensitivity of microbial respiration, nitrogen mineralization, and potential soil enzyme activities in organic alpine soils. *Glob Biogeochem Cycles* 21
- Kuo S, Sparks DL (1996) Methods of soil analysis. Part 3: Chemical methods. Soil Science Society of America, Madison, WI, 894–895
- Kwon MJ, Ballantyne A, Ciais P, Qiu C, Salmon E, Raoult N, Guenet B, Göckede M, Euskirchen ES, Nykänen H (2022) Lowering water table reduces carbon sink strength and carbon stocks in northern peatlands. *Global Change Biology* 28:6752–6770
- Lalonde K, Mucci A, Ouellet A, Gélinas Y (2012) Preservation of organic matter in sediments promoted by iron. *Nature* 483:198–200
- Lehmann J, Rillig MC, Thies J, Masiello CA, Hockaday WC, Crowley D (2011) Biochar effects on soil biota—a review. *Soil Biology and Biochemistry* 43:1812–1836

- Li Y, Yu S, Strong J, Wang H (2012) Are the biogeochemical cycles of carbon, nitrogen, sulfur, and phosphorus driven by the “Fe III–Fe II redox wheel” in dynamic redox environments? *J Soils Sediments* 12:683–693
- Liu L, Shen G, Sun M, Cao X, Shang G, Chen P (2014) Effect of biochar on nitrous oxide emission and its potential mechanisms. *J Air Waste Manag Assoc* 64:894–902
- Luo L, Meng H, Gu J-D (2017) Microbial extracellular enzymes in biogeochemical cycling of ecosystems. *Journal of Environmental Management* 197:539–549
- Luo Y, Durenkamp M, De Nobili M, Lin Q, Brookes PC (2011) Short term soil priming effects and the mineralisation of biochar following its incorporation to soils of different pH. *Soil Biology and Biochemistry* 43:2304–2316
- Mantel S, Dondeyne S, Deckers S (2023) World reference base for soil resources (WRB). *Encyclopedia of Soils in the Environment* 4:206–217
- Marsden KA, Holmberg JA, Jones DL, Charteris AF, Cárdenas LM, Chadwick DR (2019) Nitrification represents the bottle-neck of sheep urine patch N₂O emissions from extensively grazed organic soils. *Sci Total Environ* 695:133786
- Murphy J, Riley JP (1962) A modified single solution method for the determination of phosphate in natural waters. *Analytica chimica acta* 27:31–36
- Nakhavali M, Lauerwald R, Regnier P, Guenet B, Chadburn S, Friedlingstein P (2020) Leaching of dissolved organic carbon from mineral soils plays a significant role in the terrestrial carbon balance. *Glob Chang Biol* 27:1083–1096
- Oksanen J, Blanchet FG, Kindt R, Legendre P, Minchin PR, O’hara RB, Simpson GL, Solymos P, Stevens MHH, Wagner H (2013) Package ‘vegan’. *Community ecology package*, version 2, 1–295
- Op den Camp HJM, Islam T, Stott MB, Harhangi HR, Hynes A, Schouten S, Jetten MSM, Birkeland NK, Pol A, Dunfield PF (2009) Environmental, genomic and taxonomic perspectives on methanotrophic *Verrucomicrobia*. *Environ Microbiol Rep* 1:293–306
- Ozuolmez D, Na H, Lever MA, Kjeldsen KU, Jørgensen BB, Plugge CM (2015) Methanogenic archaea and sulfate reducing bacteria co-cultured on acetate: teamwork or coexistence? *Front Microbiol* 6:492
- Pandit NR, Mulder J, Hale SE, Zimmerman AR, Pandit BH, Cornelissen G (2018) Multi-year double cropping biochar field trials in Nepal: Finding the optimal biochar dose through agronomic trials and cost-benefit analysis. *Science of The Total Environment* 637:1333–1341
- Pankratov TA, Dedysh SN, Zavarzin GA (2006) The leading role of actinobacteria in aerobic cellulose degradation in Sphagnum peat bogs. *Dokl Biol Sci* 410:428–430
- Peltoniemi K, Fritze H, Laiho R (2009) Response of fungal and actinobacterial communities to water-level drawdown in boreal peatland sites. *Soil Biol Biochem* 41:1902–1914
- Pester M, Knorr K-H, Friedrich MW, Wagner M, Loy A (2012) Sulfate-reducing microorganisms in wetlands—fameless actors in carbon cycling and climate change. *Front Microbiol* 3:72
- Quilliam RS, Glanville HC, Wade SC, Jones DL (2013) Life in the ‘charosphere’—Does biochar in agricultural soil provide a significant habitat for microorganisms? *Soil Biology and Biochemistry* 65:287–293
- Razavi BS, Blagodatskaya E, Kuzyakov Y (2015) Nonlinear temperature sensitivity of enzyme kinetics explains canceling effect—a case study on loamy haplic Luvisol. *Frontiers in Microbiology* 6:1126
- Richy E, Cabello-Yeves PJ, Hernandez-Coutinho F, Rodriguez-Valera F, González-Álvarez I, Gandois L, Rigal F, Lauga B (2024) How microbial communities shape peatland carbon dynamics: new insights and implications. *Soil Biol Biochem* 191:109345
- Riedel T, Zak D, Biester H, Dittmar T (2013) Iron traps terrestrially derived dissolved organic matter at redox interfaces. *Proc Natl Acad Sci* 110:10101–10105
- Romanowicz KJ, Kane ES, Potvin LR, Daniels AL, Kolka RK, Lilleskov EA (2015) Understanding drivers of peatland extracellular enzyme activity in the PEATCosm experiment: mixed evidence for enzymic latch hypothesis. *Plant Soil* 397:371–386
- Saiya-Cork KR, Sinsabaugh RL, Zak DR (2002) The effects of long term nitrogen deposition on extracellular enzyme activity in an *Acer saccharum* forest soil. *Soil Biology and Biochemistry* 34:1309–1315
- Segata N, Izard J, Waldron L, Gevers D, Miropolsky L, Garrett WS, Huttenhower C (2011) Metagenomic biomarker discovery and explanation. *Genome biology* 12:R60
- Soares M, Rousk J (2019) Microbial growth and carbon use efficiency in soil: links to fungal-bacterial dominance, SOC-quality and stoichiometry. *Soil Biol Biochem* 131:195–205
- Spencer RGM, Hernes PJ, Aufdenkampe AK, Baker A, Gulliver P, Stubbins A, Aiken GR, Dyda RY, Butler KD, Mwamba VL (2012) An initial investigation into the organic matter biogeochemistry of the Congo River. *Geochimica et Cosmochimica Acta* 84:614–627
- Stookey LL (1970) Ferrozine—a new spectrophotometric reagent for iron. *Analytical chemistry* 42:779–781
- Taft HE, Cross PA, Jones DL (2018) Efficacy of mitigation measures for reducing greenhouse gas emissions from intensively cultivated peatlands. *Soil Biology and Biochemistry* 127:10–21
- Taghizadeh-Toosi A, Clough TJ, Sherlock RR, Condon LM (2012) Biochar adsorbed ammonia is bioavailable. *Plant and Soil* 350:57–69
- Taghizadeh-Toosi A, Elsgaard L, Clough TJ, Labouriau R, Ernstsens V, Petersen SO (2019) Regulation of N₂O emissions from acid organic soil drained for agriculture. *Biogeosciences* 16:4555–4575
- Vance ED, Brookes PC, Jenkinson DS (1987) An extraction method for measuring soil microbial biomass C. *Soil Biology and Biochemistry* 19:703–707
- Wang Y, Wang H, He J-S, Feng X (2017) Iron-mediated soil carbon response to water-table decline in an alpine wetland. *Nat Commun* 8:15972
- Wang C, Li H, Sun X, Cai T (2021) Responses of soil microbial biomass and enzyme activities to natural restoration of reclaimed temperate marshes after abandonment. *Front Environ Sci* 9:701610
- Wang S, Du Y, Liu S, Pan J, Wu F, Wang Y, Wang Y, Li H, Dong Y, Wang Z (2024a) Response of C:N:P stoichiometry to long-term drainage of peatlands: evidence from plant, soil, and enzyme. *Sci Total Environ* 919:170688
- Wen Y, Zang H, Ma Q, Evans CD, Chadwick DR, Jones DL (2019) Is the ‘enzyme latch’ or ‘iron gate’ the key to protecting soil organic carbon in peatlands? *Geoderma* 349:107–113
- Xu J, Morris PJ, Liu J, Holden J (2018) PEATMAP: Refining estimates of global peatland distribution based on a meta-analysis. *CATENA* 160:134–140
- Yamada C, Kato S, Kimura S, Ishii M, Igarashi Y (2014) Reduction of Fe (III) oxides by phylogenetically and physiologically diverse thermophilic methanogens. *FEMS Microbiology Ecology* 89:637–645
- Yang S, Li L, Peng X, Zhang R, Song L (2021) Methanogen Community Dynamics and Methanogenic Function Response to Solid Waste Decomposition. *Front Microbiol* 12:743827
- Yang T, He Q, Jiang J, Sheng L, Jiang H, He C (2022) Impact of Water Table on Methane Emission Dynamics in Terrestrial Wetlands and Implications on Strategies for Wetland Management and Restoration. *Wetlands* 42:120
- Yang Y, Sun K, Liu J, Chen Y, Han L (2022) Changes in soil properties and CO₂ emissions after biochar addition: Role of pyrolysis temperature and aging. *Science of The Total Environment* 839
- Yi X, Yi K, Fang K, Gao H, Dai W, Cao L (2019) Microbial community structures and important associations between soil nutrients and the responses of specific taxa to rice-frog cultivation. *Frontiers in Microbiology* 10:1752
- Zhang A, Zhou X, Li M, Wu HJC (2017a) Impacts of biochar addition on soil dissolved organic matter characteristics in a wheat-maize rotation system in Loess Plateau of China. 186:986–993
- Zhang C, Lin Z, Que Y, Fallah N, Tayyab M, Li S, Luo J, Zhang Z, Abubakar AY, Zhang H (2021a) Straw retention efficiently improves fungal communities and functions in the fallow ecosystem. *BMC Microbiology* 21:52
- Zhang Y, Li M, Dong L, Han C, Li M, Wu H (2021b) Effects of biochar dosage on treatment performance, enzyme activity and microbial community in aerated constructed wetlands for treating low C/N domestic sewage. *Environmental Technology & Innovation* 24:101919
- Zhang X, Han X, Yu W, Wang P, Cheng WJ Po (2017b) Priming effects on labile and stable soil organic carbon decomposition: Pulse dynamics over two years. 12:e0184978
- Zhu X, Burger M, Doane TA, Horwath WR (2013) Ammonia oxidation pathways and nitrifier denitrification are significant sources of N₂O and NO under low oxygen availability. *Proceedings of the National Academy of Sciences* 110:6328–6333
- Zimmerman AR (2010) Abiotic and microbial oxidation of laboratory-produced black carbon (biochar). *Environmental Science & Technology* 44:1295–1301